Numerical Modeling of Moist Convection in Jupiter's Atmosphere

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Introduction



Introduction



- The mean structure of the cloud layer in Jupiter's atmosphere is thought to be maintained by the statistical contribution of a large number of clouds driven by internal and radiative heating/cooling over multiple cloud life cycles.
- However, the mean structure and its relationship to cloud convection has not been clarified yet.
 - The thick visible clouds prevent the vertical structure of the entire cloud layer to be observed by remote sensing.
 - Galileo probe's entry site is one of hot spots which are cloudless region.
 - Several cloud resolving models are developed, but most of the models have been used to simulate an onset and initial expanding phase of a single cloud under a simplified and arbitrary initial condition (i.e., Yair et al., 1992, 1995; Hueso and Sanchez-Lavega, 2001).

Introduction

- The mean vertical profiles of the atmosphere have been illustrated by the results obtained using one-dimensional Weidenschilling and Lewis (1973) Atreya and Romani (1985) equilibrium cloud condensation models (ECCM)
- But, Atmospheric dynamics and cloud physical processes would modify the features obtained by ECCM.



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Fig. Vertical structure of Jupiter's cloud obtained by the equilibrium cloud condensation model (Atreya et al, 1999).

Three Cloud layers!



Toward a direct simulation

- We have been developing two-dimensional numerical models of cloud convection that incorporates phase change and cloud microphysics in order to investigate the average structure of the cloud layer that is established through a large number of life cycles of convective clouds.
 - Nakajima et al. (2000) [consider H_2O only]
 - Sugiyama et al. (2009) [consider three condensible species]



Fig: The preliminary results of numerical simulation by Sugiyama *et al.* (2009). Distribution of cloud mixing ratio (left) and vertical velocity (right). The mixing ratios of H_2O ice, NH_4SH ice, and NH_3 ice are represented using red, green, and blue color tones, respectively, and that of multiple composition cloud is represented by a superposed plot of the three colors.

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Our purpose



• In order to investigate idealistic characteristics of convective motion and mean vertical structure of the cloud layer, we perform a long-time numerical simulation of a two-dimensional cloud convection model.



Model description



Numerical model



- Two-dimensional numerical fluid model based on the quasicompressible system (Klemp and Wilhelmson, 1978)
 - The system consists of the equations of motion, continuity and thermodynamic and conservation equations of condensible species.
 - Radiation transfer process: Thermal forcing given as a substitute for radiative cooling.
 - Cloud microphysics process: The parameterization schemes of Kessler (1969) that is well-used in Earth's atmospheric simulation is used.



Set-up of the experiments



- Boundary conditions
 - Horizontal boundary is cyclic. Stress free condition and

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w = 0 are given at the lower and upper boundaries.

- Temperature and mixing ratios of vapor at the lowest level are fixed.
- Initial condition
 - Random potential temperature perturbation $(\Delta \theta_{max} = 0.1 \text{ K})$ is given to seed convective motion.



Results





Results: Animation





Temporal variation



- The convective activity of the whole layer is not steady but quasi-periodic with a period of about 40~50 days.
- Overall temperature of the cloud layer synchronizes with the intermittent convective activity.
 - We will refer the time when the active cloud convection occurs as `active period' (A) and the other as `quiet period' (Q).



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Quiet period: over view (1)

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- At the beginning of this quiet period, the NH₃ clouds are distributed horizontally.
 - Vertical motion is weak $(w = \sim 5 \text{ m/s})$
- The vertical motion in the sub-cloud layer is the remnant of convective motion driven during an active period.





Quiet period: over view (2)



- As time progress, NH₄SH clouds develop.
- Mixing of different condensible gases and condensed components across the NH₃ LCL or NH₄SH LCL is weak, but occurs occasionally due to the upward or downward penetration of convective plumes.
 - We will refer the "lifting condensation level" as LCL





Quiet period: over view (3)

- Following the onset of NH_4SH clouds, H_2O clouds begin to form locally.
- Mixing of different condensible gases and condensed components across the NH₃ LCL or NH₄SH LCL is still weak.



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Characteristic of vertical motion



- NH₃ LCL acts as a kinematic and compositional boundary, as does the NH₄SH LCL to a lesser degree.
 - The horizontal-mean mixing ratio of NH₃ gas tends to be constant between NH₄SH LCL and NH₃ LCL.
 - The vertical profile of $\sqrt{w^2}$ has local minimum or inflection points at the NH3 LCL and NH4SH LCL.





Active period: over view

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- Vigorous cumulonimbus clouds develop.
- Narrow, strong, cloudy updrafts in the cloud layer (w ~ 50 m/s)
 - Wide, weak, dry downdrafts $(w \sim 10 \text{ m/s}).$
- Considerable amounts of H_2O and NH_4SH condensates are transported to the tropopause.
- Due to the vigorous vertical transport, the mixing ratio of each condensible gases is highly inhomogeneous.



Characteristic of vertical motion



- H₂O condensation level acts as a kinematic and compositional boundary.
 - The horizontal-mean mixing ratios of all condensible gases decrease with height from the H₂O condensation level.
 - The vertical profiles of $\sqrt{w^2}$ has local minimum at the H₂O condensation level.





Mean structure: condensed components



- Horizontal mean profiles averaged over several cycles including active and quiet periods are shown.
- Considerable amounts of H_2O and NH_4SH cloud particles are observed above the NH_3 condensation.



Mean structure: Condensible gases

• The mixing ratios of NH_3 and H_2S start to decrease with height not at their respective condensation levels but at the H_2O condensation level.

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- The characteristics are not the same as that of the previous thermodynamical equilibrium calculations (ECCM).
 - Vertical mixing of dry and condensible during the active periods.



Mean structure: stability

• There is a distinct maximum of N^2 (the square of buoyancy frequency, N) at the H₂O LCL, which explains why the level acts as both a compositional and a dynamical boundary.

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- caused mainly by the change of mean molecular weight
- Weaker peaks are also present at the NH_3 and NH_4SH LCLs, which seem to act as dynamic and compositional boundaries to a certain extent during the quiet periods.



Concluding Remarks



- We perform a long-term numerical simulation with fixed thermal forcing
- The characteristics
 - Active cloud convection occurs intermittently.
 - The H₂O condensation level acts as a steady kinematic and compositional boundary.
 - The mean vertical distribution of clouds and condensible volatiles are distinctly different from those predicted by one-dimensional thermodynamical equilibrium calculations.
 - Considerable amounts of H_2O and NH_4SH cloud particles are observed above the NH_3 condensation.
 - The mixing ratios of all the volatiles start to decrease with height at H_2O condensation level.

